

冀东峪耳崖金矿床蚀变绢云母 $^{40}\text{Ar}-^{39}\text{Ar}$ 年龄 及其地质意义*

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摘要 峪耳崖大型金矿床位于华北克拉通北缘燕山造山带东段, 其主成矿阶段的时间尚待准确约束。本次工作开展了主成矿阶段矿石中蚀变绢云母 $^{40}\text{Ar}-^{39}\text{Ar}$ 同位素定年研究, 获得的坪年龄 ((169.4±1.1) Ma (MSWD=0.22)), 与反等时线年龄 ((168.4±1.8) Ma (MSWD=5.5)) 一致, 小于赋矿花岗岩的年龄 (174~175 Ma) 和辉钼矿的 Re-Os 等时线年龄 (172 Ma), 限定了峪耳崖金矿主成矿过程自 172 Ma 持续至 169 Ma。峪耳崖金矿与冀东地区其他金矿床显示相似的成岩成矿特征, 均在中侏罗世蒙古-鄂霍次克洋闭合、两侧板块碰撞造山的远程影响下, 冀东地区下地壳基底物质部分熔融形成的花岗质岩浆 (可能伴有少量幔源物质混入) 及其分异出的岩浆热液萃取基底成矿物质后, 成矿流体在 NE 向构造有利空间发生沉淀富集, 形成岩浆期后热液型金矿床。

关键词 地球化学; Ar-Ar 定年; 绢云母; 峪耳崖金矿; 冀东

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$^{40}\text{Ar}-^{39}\text{Ar}$ age of altered sericite from Yuerya Au deposit in eastern Hebei Province and its geological significance

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Abstract

Located in the east of Yanshan orogen in northern North China Craton, the Yuerya gold deposit is a meso-thermal hydrothermal deposit concern. However, the age of the main Au mineralization stage has not been constrained accurately. In this paper, $^{40}\text{Ar}-^{39}\text{Ar}$ dating was carried out on the altered sericite from ores of the main mineralization stage, which yielded a plateau age of (169.4±1.1) Ma (MSWD=0.22), consistent to the inverse isochronal age of (168.4±1.8) Ma (MSWD=5.5). This Ar-Ar age is several million years later than that of the ore-bearing granite pluton (174~175 Ma) and initial time of mineralization (172 Ma) from molybdenite Re-Os dating, suggesting that the main Au mineralization process occurred in the range of 172~169 Ma, closely related to the magmatism. The Yuerya deposit shares similar characteristics of the ore-forming granite and mineralization with other gold deposits in eastern Hebei. The remote effect of the subduction and closure of the Mongolia-Okhotsk

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Ocean and the collision orogenesis gave rise to partial melting of the lower crust in eastern Hebei, with the mixture of minor depleted mantle-derived materials. The ore-forming materials like Au and S were leached to granitic magma and the exsolved hydrothermal fluid from the basement, and then precipitated in the NE-trending fractures, resulting in the formation of magmatic hydrothermal gold deposits like the Yuerya deposit in eastern Hebei.

Key words: geochemistry, Ar-Ar dating; sericite, Yuerya gold deposit, eastern Hebei Province

峪耳崖金矿位于冀东地区宽城县境内,已查明资源量达大型规模(贾三石等,2014),是华北地区重要的金矿床之一。峪耳崖金矿的地质工作始于20世纪50年代,研究程度较高,众多学者对该矿床的区域地质、矿床地质、地球化学特征、控矿条件、矿床成因、找矿方向等方面进行了研究,取得了许多研究进展。在成矿时代方面,余昌涛等(1989)测得含金石英脉中绢云母K-Ar年龄为187.8 Ma,但K-Ar定年方法可靠程度差,且也明显大于赋矿岩体的年龄(SHRIMP 锆石U-Pb年龄为174~175 Ma,罗振宽等,2001),与地质事实不符;汤云晖等(2003)采用锆石、磷灰石裂变径迹方法获得成矿时代存在3个成矿高峰期(180~200 Ma、140 Ma和110 Ma),但该方法对热事件特别灵敏,180~200 Ma可能是成岩作用的记录,140 Ma和110 Ma是冷却年龄,与地质事实不相符;李颖等(1999)测定了峪耳崖金矿石英流体包裹体Rb-Sr等时线年龄为163.8 Ma;杨爱雪(2017)根据切穿矿体的闪长岩脉年龄间接限定峪耳崖金矿床成矿时代早于155 Ma。陈绍聪等(2014)曾对该矿床成矿早阶段辉钼矿开展了Re-Os同位素定年,获得等时线年龄为(171.9±2.7)Ma的成矿年龄。但辉钼矿Re-Os同位素体系封闭温度较高(>500℃,Suzuki et al., 1996),反映的是钼矿化的年龄,金矿是钼矿成矿之后形成的,成矿温度介于118~410℃之间(邱检生等,1994;贾三石等,2014;郭博巍等,2017),因此,辉钼矿Re-Os年龄限定了峪耳崖金矿成矿时代的最大值,没有确定主成矿阶段的准确时间,不利于深入理解该矿床的金矿成矿作用过程。

本次工作在野外地质调查研究的基础上,对峪耳崖主成矿阶段矿石中蚀变矿物绢云母开展⁴⁰Ar-³⁹Ar精细测年工作,厘定了主成矿阶段的时限,并结合区域成岩成矿信息,探讨了冀东地区金矿床及相关花岗岩的成岩成矿演化过程。

1 区域地质

冀东地区位于华北克拉通北缘燕山造山带的东

段(图1),地质演化历史长,地质构造复杂,岩浆活动强烈,矿产资源丰富。区内构造主要发育宽缓的褶皱和断裂,区域性褶皱构造有近EW向展布的马兰峪复式背斜,其核部出露太古宙结晶基底,岩性以辉石麻粒岩、黑云斜长片麻岩为主,两翼主要为中-新元古界的碎屑岩、黏土岩和碳酸盐岩。区域断裂带多呈NE向、NNE向和近EW向,如近EW向喜峰口-青龙断裂带、NE向凌源-喜峰口断裂带、NW向冷口断裂带,它们控制了区域中酸性岩浆岩的侵入和空间展布(肖振等,2010),对区内金多金属成矿作用具有重要意义。

根据成岩时代,冀东地区显生宙岩浆活动主要可划分为海西期、印支期和燕山期3期。海西期岩浆岩在区内出露有限,主要为晚石炭世的东湾子超基性岩(Zhao et al., 2007;贺文等,2015)。印支期岩浆岩以花岗岩为主,主要出露于冀东中部,虽发育较少,但形成了区内规模最大的都山复式花岗岩体,时代属晚三叠世(罗振宽等,2003;Jiang et al., 2018)。燕山期花岗岩类发育最为广泛,花岗岩带呈近EW向展布,单个岩体长轴方向通常为NE向,如早侏罗世青山口、高家店、蛇盘兔岩体等(罗振宽等,2001;尹业长等,2018;Jiang et al., 2018),中侏罗世峪耳崖、牛心山、唐杖子、金宝沟等花岗(斑)岩体(罗振宽等,2001;贺文,2015;李曼等,2016);燕山晚期(晚侏罗世—早白垩世)的岩浆岩主要产于兴隆-平泉-凌源一线西北侧(熊乐,2017)及北京密云附近。

冀东地区具有优越的成矿地质条件,区内发育许多金、钼(铜)矿床(图1),主要包括金厂峪、峪耳崖、牛心山、唐杖子等大、中型金矿床等。金矿床总体沿马兰峪复式背斜展布。区内显生宙的成矿事件主要发生于三叠纪、中侏罗世、白垩纪3个时期。金厂峪金矿床的成矿时代虽长期存在争议,但其辉钼矿Re-Os年龄为(243±7)Ma(宋扬等,2011),可能代表了冀东地区三叠纪的成矿作用。侏罗纪的金、钼成矿作用在该区非常发育,如中侏罗世的峪耳崖、唐杖子、牛心山等金矿床(白洪生,1992;李正远等,

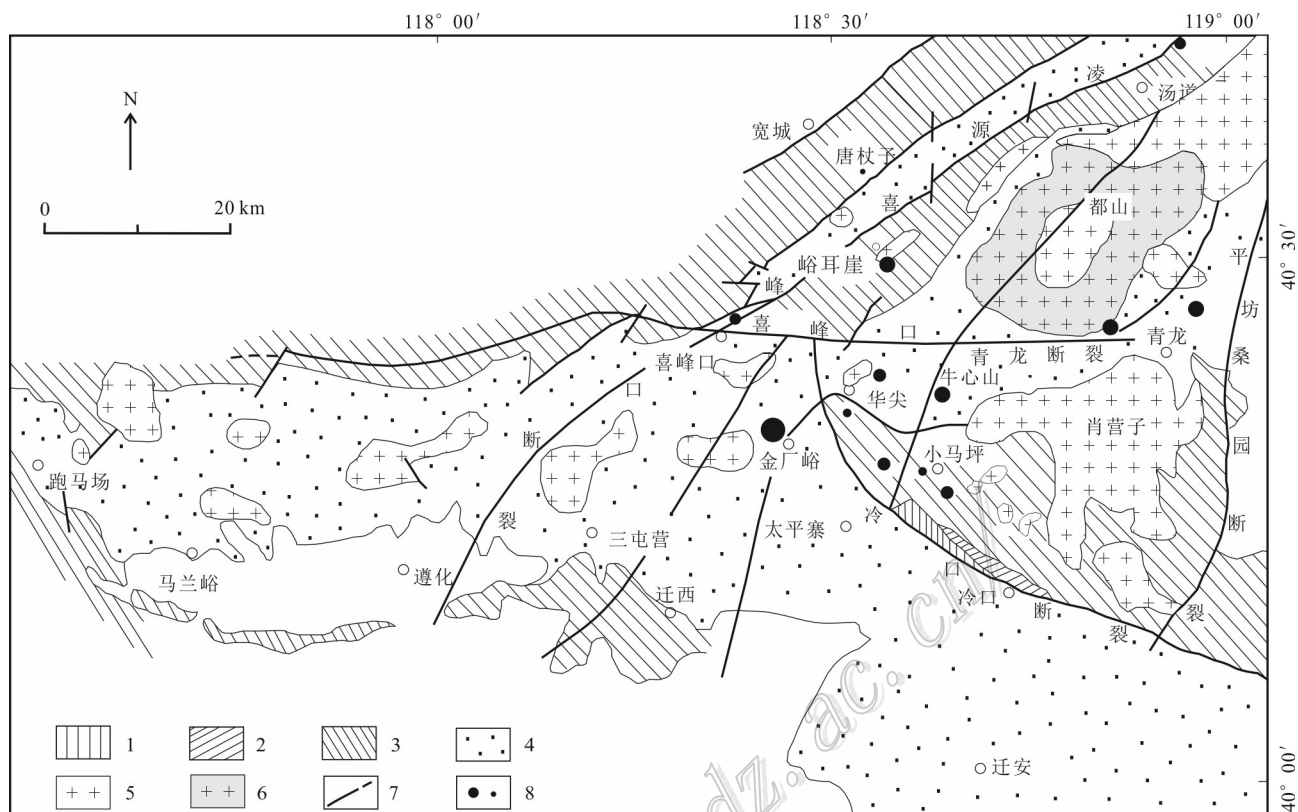


图1 冀东地区区域地质及金矿床分布简图(据李俊健等,2002修改)

1—中生界;2—古生界;3—中-新元古界;4—太古宙变质杂岩;5—燕山期花岗岩;6—印支期花岗岩;7—断层;8—金矿床(点)

Fig. 1 Sketch map of regional geology and ore distribution in the eastern Hebei (modified after Li et al., 2002)

1—Mesozoic; 2—Paleozoic; 3—Meso-Neoproterozoic; 4—Archean metamorphic rock; 5—Yanshanian granite; 6—Indosinian granite; 7—Fault; 8—Gold deposit (ore spot)

2014;陈绍聪等,2014)和太平村钼矿(孙金龙,2017),金矿类型有蚀变岩型、石英脉型、隐爆角砾岩型。晚侏罗世—早白垩世的矿化以铜、钼为主,如小寺沟铜钼矿床、寿王坟铜矿床(张瑞斌等,2008)。

2 矿床地质

峪耳崖金矿区位于马兰峪复背斜的北翼,出露地层主要为中元古界长城系高于庄组碳酸盐岩,岩性以白云质灰岩和灰质白云岩为主(图2)。

矿区出露的岩浆岩有峪耳崖花岗岩体,其呈NE向侵位于中元古界长城系高于庄组碳酸盐岩地层中。该岩体由灰白色中细粒黑云母花岗岩和肉红色中粗粒黑云母花岗岩组成,前者多分布于岩体的顶部、边部,发育白云岩围岩捕掳体、顶垂体;后者多分布于岩体中、深部。2类花岗岩均呈块状构造,主要矿物有钾长石(50%~55%)、斜长石(15%)、石英

(25%~30%),可见少量黑云母,副矿物有磷灰石、锆石、黄铁矿等。在结构和钾长石类型上,2种岩性存在一定差异:灰白色花岗岩呈中粒-似斑状结构,钾长石以微斜长石为主;肉红色花岗岩为粗粒花岗结构,钾长石以正长石为主。2种花岗岩之间无明显的界线。灰白色花岗岩和肉红色花岗岩形成时代分别为(175±1)Ma和(174±3)Ma(罗振宽等,2001),属同一期次不同阶段的产物,成岩物质可能主要来自太古代基底的部分熔融(罗振宽等,2001;陈冬,2007)。此外,矿区岩脉发育,主要有黑云母石英闪长岩、黑云角闪闪长岩和角闪闪长岩,形成于163~155 Ma(杨爱雪,2017)。

断裂构造发育,主要有NE向、NNE向、NEE向、NW向、近EW向5组。其中,NE向断裂带在矿区内分布广泛,倾向NW,倾角一般较缓,为30°~45°,北部较陡,为60°~75°,断裂带内发育片理化透镜体,属逆断层,局部显示压扭性特征,控制着矿区65%左右

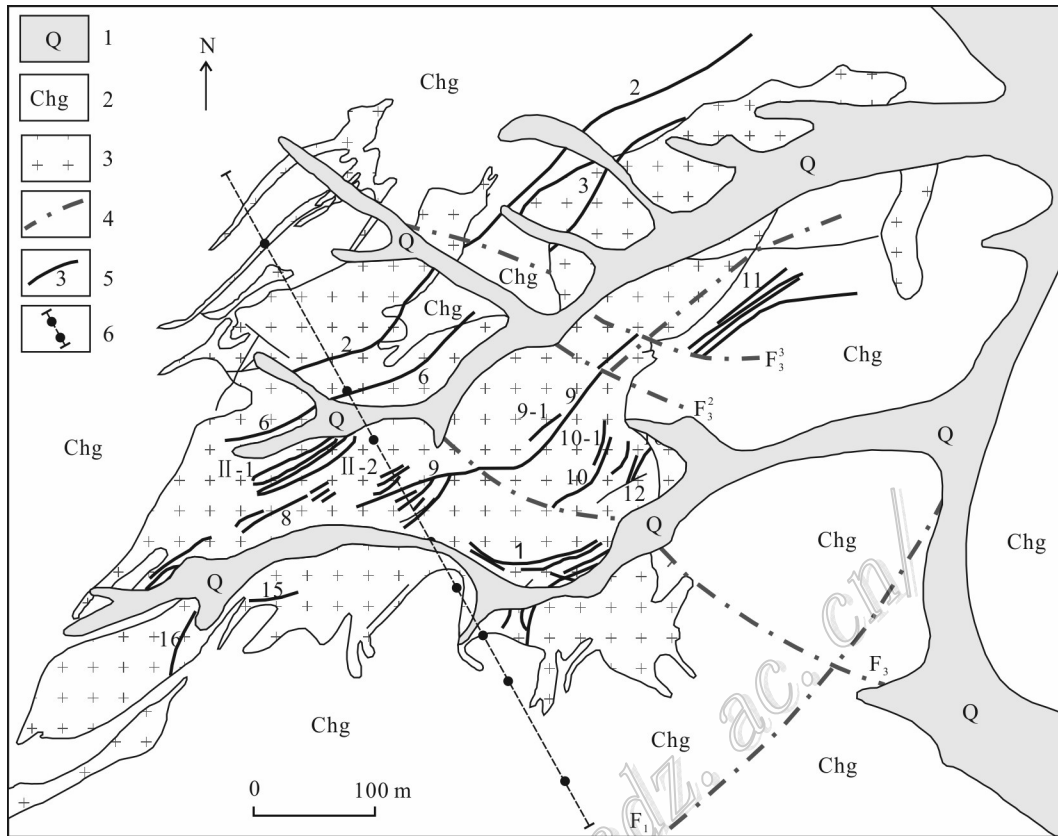


图2 峪耳崖金矿床地质图(据肖振等,2009修改)

1—第四系;2—高于庄组灰质白云岩;3—花岗岩;4—断裂;5—矿脉及编号;6—2号勘探线及钻孔位置

Fig. 2 Geological map of the Yuerya gold deposit (modified after Xiao et al., 2009)

1—Quaternary; 2—Gaoyuzhuang Formation grey dolomite; 3—Granite; 4—Fault; 5—Orebody and its serial number; 6—No. 2 exploration line and drill hole

的金矿脉产出(金天尚,2003)。NNE向断裂带主要分布于矿区西南部位,倾向北西,倾角 $37^{\circ}\sim 66^{\circ}$,发育碎裂岩、片理化、斜擦痕,显示压扭性特征。相比NE向断裂带,NEE向断裂带发育相对较弱,局部密集,以压性为主,也被矿脉充填。近EW向断裂在矿区南部略有发育,显示压扭性特征,部分充填含金石英脉。NW向断裂带仅局部可见,属张性断裂,在成矿早期控制了少部分矿脉,但更主要的是切割NE向矿体,是成矿后断裂,对矿体起破坏作用。

金矿多集中分布于岩体南部、中部和北东部,可分为南矿带、中矿带和北矿带。金矿脉严格受NE向、NNE向、NEE向断裂带控制。矿山勘探资料显示,金矿体多为盲矿体,赋存于花岗岩体内(图2,图3),少数分布于接触带围岩中的断裂带内(封文学等,2017)。矿化类型可以分为含金石英脉型和微细脉浸染状蚀变岩型,前者主要产于花岗岩体及其外

接触带灰质白云岩的断裂破碎带中,石英脉厚度大于5 cm;呈乳白色者所含硫化物较少,金品位低,达不到工业要求;呈烟灰色者,发育团块状、浸染状黄铁矿、黄铜矿、方铅矿等多金属硫化物,金品位较高,构成重要的金矿体。后者分布于含矿断裂带两侧的次级裂隙中,表现为含金石英细脉呈树枝状、网脉状,单脉厚度小于1 cm,黄铁矿呈浸染状、团块状产于细脉及两侧花岗岩中,围岩蚀变强烈,主要为黄铁矿化、绢云母化、硅化等。

金矿石矿物主要有黄铁矿、黄铜矿,其次为磁黄铁矿、闪锌矿、辉铜矿,少量的辉钼矿,微量的自然金、自然铋等;脉石矿物以石英为主,还可见绢云母、绿泥石、方解石、铁白云石等。金通常呈显微金产于矿石中,局部品位高的矿石中可见自然金。金主要以包裹金、裂隙金和晶隙金赋存于黄铁矿、磁黄铁矿等硫化物的晶格缺陷、裂隙、晶隙中。部分金以

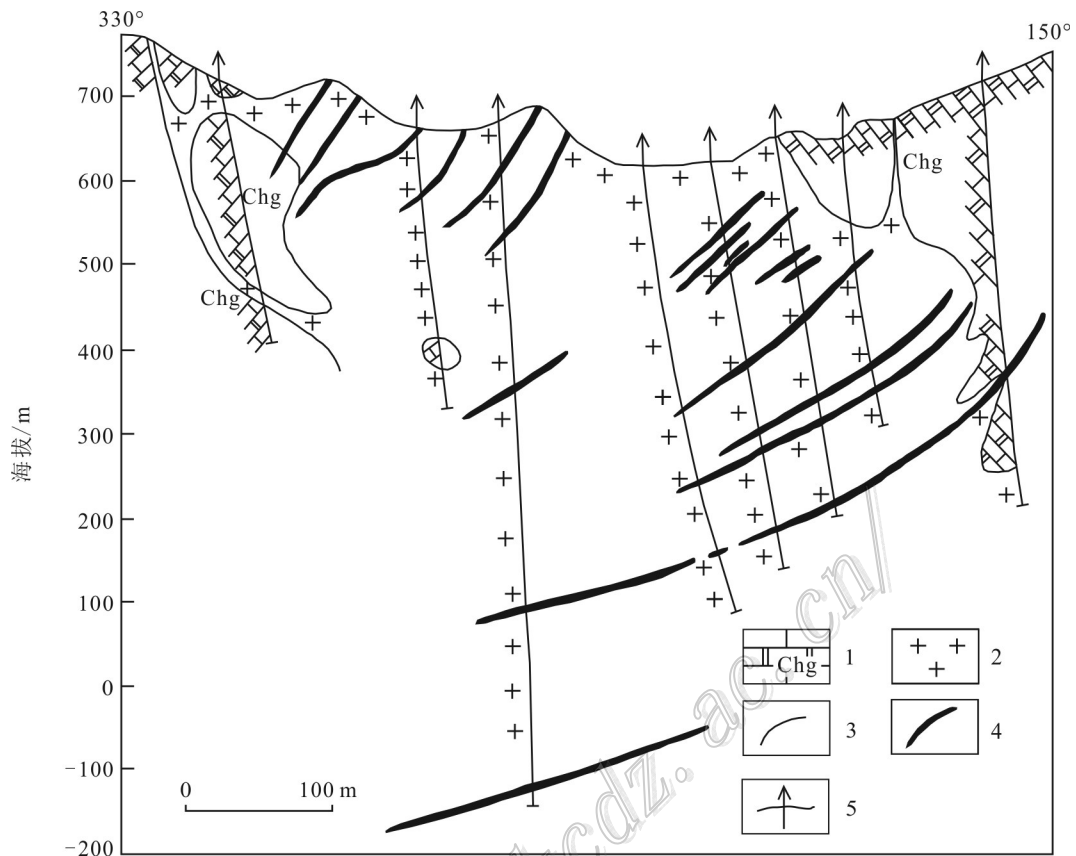


图 3 峪耳崖金矿 2 号勘探线剖面图(据肖振等,2010)

1—灰质白云岩;2—花岗岩;3—地质界线;4—矿脉;5—钻孔

Fig. 3 Geological section along No. 2 exploration line of the Yuerya gold deposit (after Xiao et al., 2010)

1—Lime dolomite; 2—Granite; 3—Geological boundary; 4—Orebody; 5—Drill hole

薄膜金的形式吸附于褐铁矿表面(封文学,2017)。

石英脉型金矿石具脉状构造,粒状结晶结构、碎裂结构,黄铁矿、黄铜矿、方铅矿、闪锌矿等硫化物多呈浸染状、脉状、细脉状、条带状、角砾状等构造,自形-半自形及他形细粒结晶结构、包含结构、碎裂结构、交代结构和充填结构等。

围岩蚀变强烈,以黄铁矿化、绢云母化、硅化、黄铁绢英岩化、钾化为主,其次为绿泥石化、高岭土化等。其中,黄铁矿化、硅化、绢云母化与矿化关系最为密切。花岗岩体中以硅化、钾化、黄铁绢英岩化为主;在岩体与碳酸盐岩接触的外带,蚀变以强烈硅化、钾化为主,局部发育矽卡岩化、大理岩化。

根据矿物共生组合及相互穿插关系,可将热液成矿作用划分为 5 个阶段。黄铁矿-石英阶段(I),形成乳白色星点状粗粒自形-半自形黄铁矿石英脉,围岩蚀变以硅化为主。石英-辉钼矿阶段(II),形成

辉钼矿-石英脉,石英呈烟灰色,黄铁矿多呈半自形-他形,辉钼矿有 2 种:一种表现为烟灰色石英脉边部的薄层状集合体,辉钼矿呈弯曲叶片状,粒度较小;另一种表现为附于石英脉脉壁的薄膜状集合体,辉钼矿呈叶片状,粒度更细。此外,还发育磁黄铁矿、黄铜矿等,偶见自然金,围岩蚀变主要有硅化、黄铁矿化、绿帘石化、铁白云石化。石英-黄铁矿阶段(III),形成黄铁矿-石英脉,石英多呈烟灰色,半自形-他形黄铁矿呈细脉或网脉状充填裂隙,脉两侧花岗岩黄铁绢英岩化强烈,该阶段矿石金品位高,是主要的金成矿阶段。石英-多金属硫化物阶段(IV),形成方铅矿-闪锌矿-黄铁矿-石英微细脉,发育自然金,含银较高(陈冬,2007),围岩蚀变有硅化、黄铁矿化、弱绿泥石化、弱碳酸盐化,是重要的成金阶段。碳酸盐阶段(V),形成方解石脉及少量石英脉,含少量黄铁矿,但基本不含金,围岩蚀变以绿泥石化、碳酸盐化为主。

3 样品制备与测试结果

3.1 样品特征

本次工作用于定年的样品(YEY-B3)采自含金石英脉两侧的蚀变灰白色花岗岩(图4a),属主成矿阶段(Ⅲ阶段)矿石,采样位置地理坐标为(E118°32'17";N40°29'15")。该矿石产于白色花岗岩与高于庄组白云质灰岩接触带附近,矿石呈块状构造、半自形粒状结构,金属矿物主要为半自形-他形黄铁矿,脉石矿物有石英、绢云母、长石、黑云母等,围岩蚀变以黄铁矿化、绢云母化、硅化为主,还发育弱绿帘石化、绿泥石化。其中,绢云母呈鳞片状集合体,产于石英脉与花岗岩接触部位(图4c),与黄铁矿共生,是含金热液蚀变的产物;同时花岗岩所含斜长石发生绢云母化,但该类型绢云母较前者更细小(图4b、f)。本次工作的测试对象即为产于含金石英脉与花岗岩接触部位集合体中的绢云母。

3.2 测试方法

绢云母单矿物的分离是在河北省廊坊市诚信地质服务有限公司完成。矿石样品经粉碎、分离、粗选和精选,获得纯度>99%的绢云母单矿物样品,再经超声波清洗。清洗后的样品被封进石英瓶中送核反应堆中接受中子照射。照射工作是在中国原子能科学研究院的“游泳池堆”中进行的,使用B4孔道,中子流密度约为 $2.60 \times 10^{13} \text{ n cm}^{-2} \text{ s}^{-1}$ 。照射总时间为1444 min,积分中子通量为 $2.25 \times 10^{18} \text{ n cm}^{-2}$;同期接受中子照射的还有用做监控样的标准样:ZBH-25黑云母标样,其标准年龄为 $(132.7 \pm 1.2) \text{ Ma}$, $w(\text{K})$ 为7.6%。

样品的阶段升温加热使用石墨炉,每一个阶段加热30 min,净化30 min。质谱分析在中国地质科学院地质研究所Ar-Ar年代学同位素实验室多接收稀有气体质谱仪Helix MC上进行,每个峰值均采集20组数据。所有的数据在回归到时间零点值后再进行质量歧视校正、大气氩校正、空白校正和干扰元素同位素校正。中子照射过程中所产生的干扰同位素校正系数通过分析照射过的 K_2SO_4 和 CaF_2 来获得,其值为: $(^{36}\text{Ar}/^{37}\text{Ar}_0)_{\text{Ca}}=0.000\ 2389$, $(^{40}\text{Ar}/^{39}\text{Ar})_{\text{K}}=0.004\ 782$, $(^{39}\text{Ar}/^{37}\text{Ar}_0)_{\text{Ca}}=0.000\ 806$ 。 ^{37}Ar 经过放射性衰变校正; ^{40}K 衰变常数 $\lambda=5.543 \times 10^{-10} \text{ a}^{-1}$;用ISO-PLOT程序计算坪年龄及正、反等时线(Ludwig, 2001)。坪年龄误差以 2σ 给出。详细实验流程见有关文献(陈文等,2006;张彦等,2006)。

3.3 测试结果

峪耳崖金矿矿化蚀变绢云母 ^{40}Ar - ^{39}Ar 同位素定年分析结果见表1。

在700~1400℃之间共计13个温度阶段的加热结果组成一个基本没有扰动的年龄图谱(图5),全年龄为168.9 Ma。其中,1000~1280℃的9个阶段视年龄之间的差异极小,计算其坪年龄为 $(169.4 \pm 1.1) \text{ Ma}$ (2σ)(MSWD=0.22),对应的 ^{39}Ar 析出量为94.9%。反等时线年龄为 $(168.4 \pm 1.8) \text{ Ma}$ (MSWD=5.5), $^{40}\text{Ar}/^{36}\text{Ar}$ 初始值为 (364 ± 63) ,反等时线年龄和坪年龄在误差范围内完全一致(图5)。

4 讨论

4.1 成矿时代

蚀变成因的绢云母常在各类金属热液矿床中作为脉石矿物出现,被广泛用于测定Ar-Ar同位素年龄,从而确定成矿时代。绢云母Ar-Ar同位素体系对后期地质作用十分敏感,容易遭受后期地质作用叠加改造,出现明显的扩散丢失图谱(Hanson et al., 1975),而未受扰动的绢云母则形成平坦型年龄图谱(邱华宁等,1997;陈文等,2011)。本次测定的绢云母呈平坦型年龄谱,未出现异常的年龄谱,表明绢云母在169 Ma左右形成之后没有再受到高于其封闭温度的构造-热事件的影响。因此,本次所获得绢云母坪年龄 $((169.4 \pm 1.1) \text{ Ma})$,MSWD=0.22)可以代表其形成的冷却年龄。

前人对峪耳崖金矿床成矿流体的测温结果显示高度的一致性,均一温度介于118~395℃(邱检生等,1994;贾三石等,2014;郭博巍等,2017)。该矿床内辉钼矿形成于第Ⅱ成矿阶段,辉钼矿结晶沉淀温度明显低于其Re-Os同位素体系封闭温度($>500^\circ\text{C}$, Suzuki et al., 1996),说明辉钼矿于成矿早阶段沉淀后,Re-Os体系即封闭,故而先前工作测得的171.9 Ma的辉钼矿Re-Os年龄代表了主成矿阶段初始的时间(陈绍聪等,2014)。而本次工作所测试绢云母样品采自第Ⅲ成矿阶段的金矿石,主成矿阶段成矿温度与绢云母Ar同位素体系封闭温度(绢云母Ar同位素体系封闭温度 $(350 \pm 50)^\circ\text{C}$,邱华宁等,1997;陈文等,2006)基本一致,表明与金矿同时形成的绢云母在主成矿阶段达到封闭,因此绢云母Ar-Ar冷却年龄记录了峪耳崖金矿床主成矿阶段延续至169 Ma。而155 Ma的角闪闪长岩脉切穿了金矿体(杨爱雪,

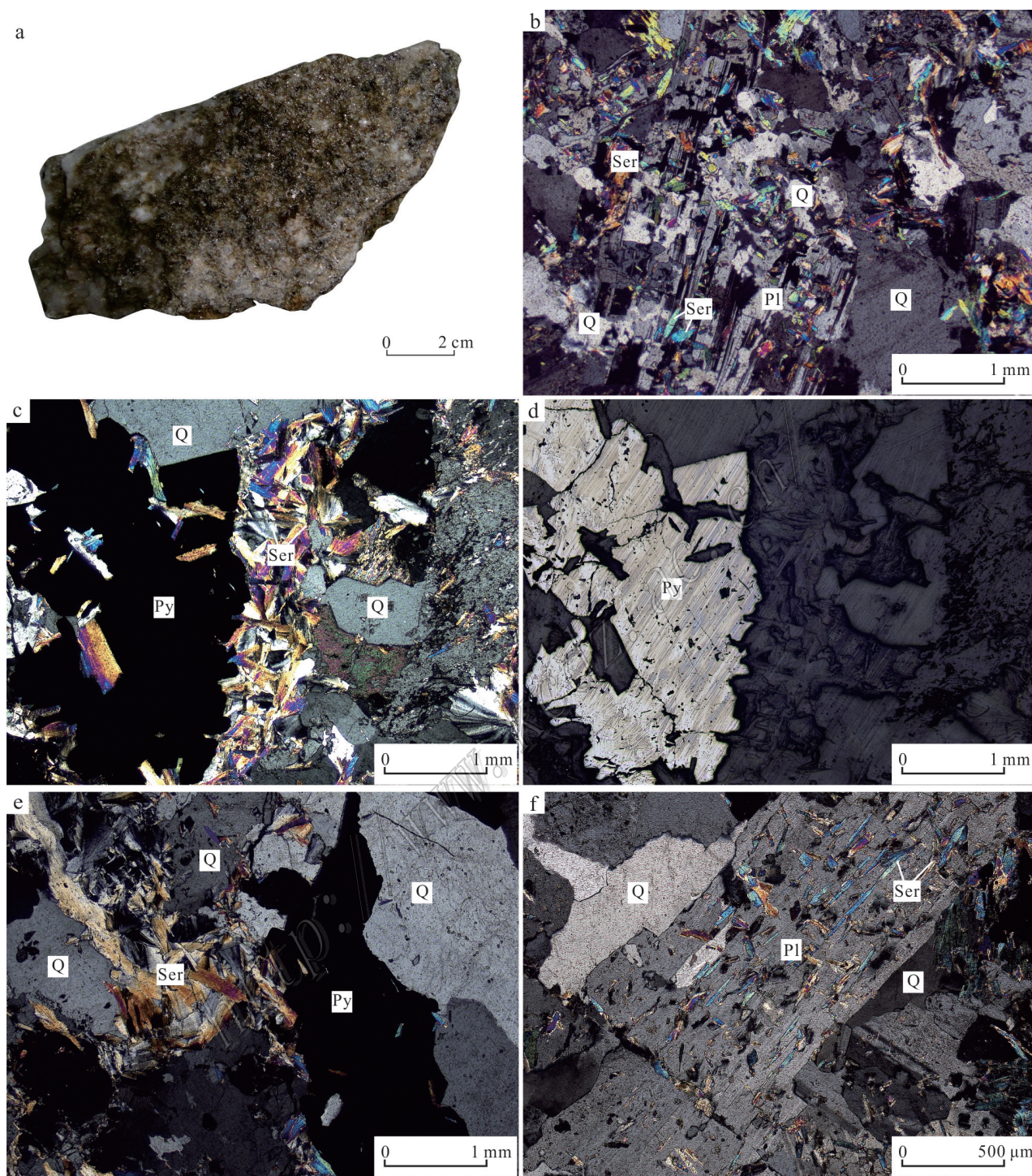


图 4 峪耳崖金矿床绢云母样品特征

a. 黄铁绢英岩化花岗岩; b. 花岗岩斜长石强烈绢云母化; c, d. 黄铁绢英岩化, 绢云母呈鳞片状集合体产于石英脉与花岗岩接触部位 (分别为透射光和反射光图像); e. 黄铁绢英岩化; f. 斜长石发生绢云母化蚀变

Q—石英; Pl—斜长石; Ser—绢云母; Py—黄铁矿

Fig. 4 Characteristics of the sericite sample from the Yuerya gold deposit

a. Sericite-bearing granite; b. Intense sericitization of plagioclase in granite; c, d. Beresitized ore, and sericite with pyrite in the contact zone between the quartz vein and the granite under transmission light and reflected light, respectively; e. Beresitized ore; f. Sericitization of plagioclase

Q—Quartz; Pl—Plagioclase; Ser—Sericite; Py—Pyrite

表1 峪耳崖金矿床绢云母⁴⁰Ar-³⁹Ar同位素分析结果

Table 1 The ⁴⁰Ar-³⁹Ar isotopic analyses of the sericite from the Yuerya gold deposit

θ/\square	$(^{40}\text{Ar}/^{39}\text{Ar})_m$	$(^{36}\text{Ar}/^{39}\text{Ar})_m$	$(^{37}\text{Ar}/^{39}\text{Ar})_m$	$(^{38}\text{Ar}/^{39}\text{Ar})_m$	$^{40}\text{Ar}/\%$	F	$^{39}\text{Ar}/10^{-14}\text{ mol}$	^{39}Ar 积累/ $\%$	t/Ma	$1\sigma/\text{Ma}$
700	31.2876	0.0311	0	0.0135	70.63	22.0999	0.05	0.10	179	16
800	23.3501	0.0329	0	0.0000	58.39	13.6330	0.02	0.13	112	50
900	21.6225	0.0063	0	0.0142	91.30	19.7424	2.22	4.65	160.5	1.6
1000	22.2135	0.0036	0.0001	0.0136	95.21	21.1486	3.73	12.21	171.4	1.7
1070	21.6412	0.0027	0	0.0133	96.34	20.8502	5.00	22.34	169.1	1.6
1090	21.0714	0.0009	0.0001	0.0129	98.72	20.8007	8.30	39.17	168.7	1.6
1110	21.0835	0.0009	0	0.0129	98.77	20.8234	9.35	58.13	168.9	1.6
1130	21.2094	0.0012	0	0.0131	98.30	20.8492	7.51	73.36	169.1	1.6
1150	21.2416	0.0011	0	0.0131	98.40	20.9019	5.04	83.57	169.5	1.6
1180	21.2412	0.0012	0	0.0129	98.29	20.8799	4.25	92.20	169.4	1.6
1220	21.1473	0.0009	0.0037	0.0132	98.75	20.8824	2.41	97.08	169.4	1.7
1280	21.3159	0.0014	0	0.0129	97.99	20.8865	1.24	99.59	169.4	1.8
1400	25.2745	0.0108	0	0.0149	87.41	22.0921	0.20	100.00	178.7	3.3

注:下标m代表样品中测定的同位素比值;全年龄=168.9 Ma;样品质量=27.67 mg;照射参数J=0.004714。

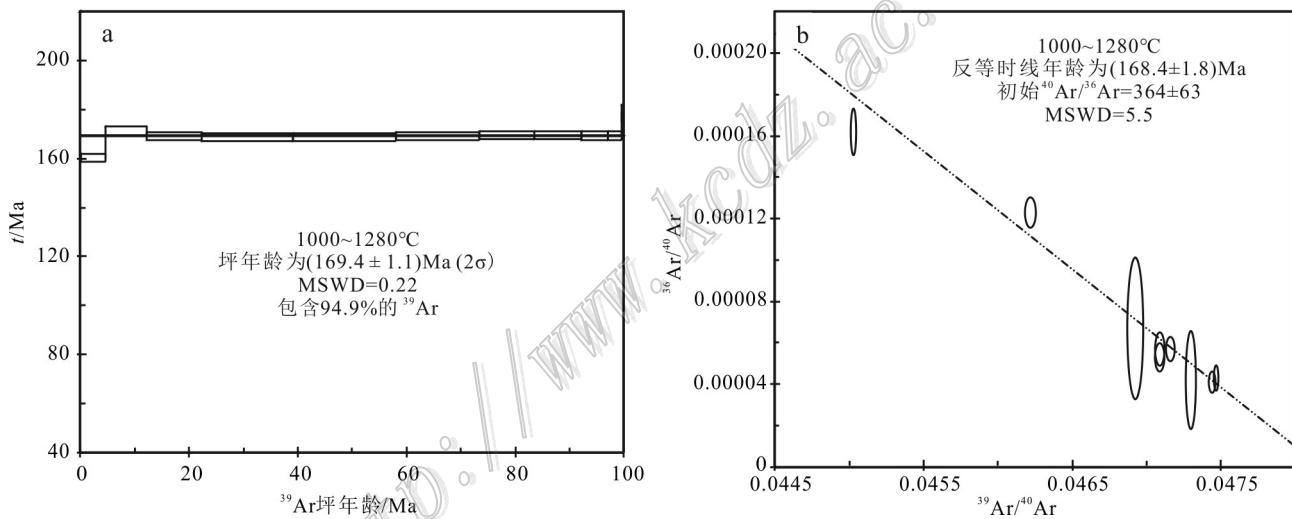


图5 峪耳崖金矿床热液蚀变绢云母的⁴⁰Ar-³⁹Ar坪年龄(a)和反等时线年龄图(b)

Fig5 ⁴⁰Ar-³⁹Ar spectrum (a) and inverse isochronal age (b) diagrams of the altered sericite from the Yuerya gold deposit

2017),表明此时峪耳崖金成矿作用已经结束。

4.2 区域成岩成矿特征

冀东地区中生代花岗质岩浆作用可以划分为5期(戴雪灵等,2010;李强等,2012;杨付领等,2015;徐希阳等,2016;Jiang et al., 2018):晚三叠世(223~206 Ma),以都山花岗岩基为代表;早侏罗世(199~187 Ma),以青山口、罗文峪、蛇盘兔和肖营子花岗岩体为代表;中侏罗世早期(177~168 Ma),以牛心山、峪耳崖、白家店、达子沟花岗岩体为代表;晚侏罗世(约160 Ma),以七拨子岩体为代表;早白垩世(122

Ma),以小寺沟岩体、寿王坟斑岩体为代表。其中,中侏罗世早期花岗岩类一般属高钾钙碱性系列,准铝质-弱过铝质,显示高Si、碱,低Mg、Fe、Ti特征, $w(\text{Al}_2\text{O}_3)$ 介于12%~16%;富集大离子亲石元素、亏损高场强元素,低Sr、Y、Yb含量;稀土元素总含量低,富集轻稀土元素,稀土元素配分模式呈右倾型,弱至中等负Eu异常,显示不同程度分异特征;全岩Sr同位素初始比值(I_{Sr})介于0.7051~0.7069, ϵ_{Nd} 介于-10.8~-15.3,锆石 ϵ_{Nd} 介于-8.2~-24.4(康显桂等,1998;贺文,2015;李曼等,2016;Liu et al., 2016;徐希

阳等,2016;封文学,2017;尹业长等,2018)。李承东等(2004)将这类花岗岩划分为低Sr、低Y型花岗岩,区别于冀东地区中晚三叠世一早侏罗世高Sr、低Y花岗岩类。峪耳崖赋矿花岗岩分异指数较未成矿岩体更高,负Eu异常更显著(封文学,2017),推测区域金成矿作用与花岗质岩浆的高分异作用存在一定关系。

区内牛心山金矿产于花岗岩内及接触带古太古界迁西群裂隙内,唐杖子金矿主要产于花岗斑岩体与隐爆角砾岩体内断裂带中,二者矿体呈透镜状、似层状、脉状,矿脉走向主体为NE向。矿化类型以蚀变岩型、石英脉型、细脉浸染型为主,还有部分蚀变破碎带型、隐爆角砾岩型等。矿石一般呈浸染状、团块状、脉状、细脉状构造,他形、半自形-自形粒状、交代结构等。矿石矿物包括黄铁矿、磁黄铁矿、黄铜矿、方铅矿、闪锌矿、黝铜矿、辉钼矿、自然金等,脉石矿物有石英、钾长石、斜长石、绢云母、绿泥石、绿帘石、方解石等。围岩蚀变可见黄铁矿化、硅化、绢云母化、黄铁绢英岩化、绿帘石化、绿泥石化、碳酸盐化等(舒航等,1989;贺文,2015)。成矿流体H、O、He、C、S等同位素组成变化范围相对较窄,主体为岩浆水,显示中温、低盐度、低密度的特征(贺文,2015;石成龙等,2015)。前人研究认为其成矿物质一般来自岩浆及区域基底地层变质岩系(贺文,2015;Kong et al., 2015; Liu et al., 2016)。成矿时代为169~176 Ma(白洪生,1992;李正远等,2014),在误差范围内与赋矿岩体时代一致。总之,在成矿地质特征、赋矿岩体、成岩成矿时代、成矿流体和物质来源等方面,这2个金矿床均与峪耳崖金矿床表现出高度的相似性,表明它们应该经历了类似的成矿作用过程。

4.3 成岩成矿演化过程

受西伯利亚板块、古亚洲洋、蒙古-鄂霍茨克洋联合块体和华北板块相互作用的影响,华北克拉通北缘发生了多期次强烈的构造-岩浆活动(聂凤军等,2011;Li et al., 2018; Wang et al., 2018)。晚二叠世—三叠纪初期,华北板块与西伯利亚板块发生碰撞(赵越等,2010;Wilde, 2015; Li et al., 2017),晚三叠世碰撞后岩石圈伸展体制下形成的花岗岩类标志着古亚洲洋构造演化的结束(叶浩,2014;Liu et al., 2018)。华北克拉通北缘发育大量燕山期岩浆-金属成矿作用,可能与蒙古-鄂霍茨克洋和古太平洋板块的俯冲、叠合、转换以及大洋消减-闭合后的伸展等动力学过程有密切联系(郑建平等,2017)。古地磁

研究表明,蒙古-鄂霍茨克洋于二叠纪末期开始缓慢闭合,并且自西向东呈“剪刀式”闭合(Zorin, 1999; Kravchinsky et al., 2010; Donskaya et al., 2013)。洋盆的西部闭合较早,约在早-中侏罗世(Zorin, 1999; Donskaya et al., 2013),引起小兴安岭西北部至冀北-辽西地区的陆壳加厚和逆冲推覆事件(李宇等, 2015);而中、东部的闭合发生于晚侏罗世—早白垩世,两侧地体内发育大型逆冲推覆构造,最终形成蒙古-鄂霍茨克构造带(Cogné et al., 2005; Sun et al., 2013)。华北克拉通北缘阴山-燕山陆内造山带早、中侏罗世发生强烈挤压和碰撞,通常被认为是对蒙古-鄂霍茨克洋盆关闭的远程陆内响应,SN向构造挤压形成一系列近EW向延伸的褶皱-逆冲构造(张岳桥等,2007;王永超,2017)。冀东NE向构造可能也形成于蒙古-鄂霍茨克洋闭合所引起的南北向挤压体制之下(Davis et al., 2001; Dong et al., 2015)。毛景文等(2003;2005)指出自印支晚期—中侏罗世,华北克拉通在一定程度上仍处于碰撞造山阶段,在碰撞造山的伸展期,大量花岗岩浆经过同熔或重熔作用生成及上侵定位,引发了中国北方200~160 Ma期间的大规模成矿作用。而对于古太平洋板块向西俯冲启动的时限,大部分学者认为开始于早侏罗世末—中晚侏罗世(Bartolini et al., 2001; Davis et al., 2001; Wang et al., 2018; 郑永飞等,2018)。黄始琪等(2014)在蒙古-鄂霍茨克造山带中段(蒙古境内的阿穆尔板块)断裂构造中识别出NW-SE向挤压应力场,认为可能对应了中晚侏罗世—白垩纪古太平洋板块向西俯冲的远程影响。Wang等(2018)认为古太平洋板块俯冲作用影响燕山地区的时间为165~160 Ma。总之,古太平洋板块向西俯冲可能对冀东地区产生了重要的影响,但明显晚于冀东地区170 Ma左右的成矿事件,即冀东地区中侏罗世岩浆-成矿作用可能与蒙古-鄂霍茨克洋的俯冲和闭合碰撞关系更为紧密。

在鄂霍茨克洋闭合、两侧地块碰撞造山的远程影响下,冀东地区近EW向断裂再次活化,NE向断裂形成(陈云峰,2007;Wang et al., 2018),为岩浆及流体活动提供了通道和定位空间。由于地壳增厚升温或局部伸展降压,古老基底发生部分熔融形成花岗质岩浆,同时伴有少量亏损地幔来源的镁铁质岩浆混入(徐希阳等,2016),沿EW向区域深断裂向上侵位,同时萃取了下地壳Au等成矿物质,形成了174~175 Ma低Sr、低Y的深源花岗岩类(李承东等,

2004)。岩浆演化后期,分异出H、O、He、C、S等同位素组成变化范围相对较窄的含金热液流体(邱检生等,1994;Kong et al., 2015; Liu et al., 2016)。在地壳浅部应力松弛阶段,NE向断裂发生张性或扭张性活动,瞬时真空产生的“泵吸作用”使流体在其中沉淀形成金矿脉(宋扬等,2013),从而形成冀东峪耳崖等169~176 Ma的岩浆热液型金矿床。此外,杨爱雪(2016)发现峪耳崖矿区部分163 Ma沿矿脉平行产出的NE向闪长岩脉亦发育金矿化,但这是早期金矿化的延续,还是幔源岩浆(杨爱雪,2016)所带来的后期矿化叠加,还需要更多的工作来证实。而155 Ma的闪长岩脉切穿金矿体,表明峪耳崖金成矿作用此时已经结束。

5 结论

(1) 冀东峪耳崖金矿主成矿阶段蚀变绢云母 ^{40}Ar - ^{39}Ar 坪年龄为 $(169.4\pm 1.1)\text{Ma}$ (MSWD=0.22)。结合前人资料,认为峪耳崖金矿主成矿阶段始于172 Ma,延续至169 Ma,155 Ma时成矿作用已经结束。

(2) 峪耳崖金矿与冀东地区其他岩浆热液型金矿床,均是在中侏罗世蒙古-鄂霍次克洋闭合、两侧板块碰撞造山的远程影响下,基底部分熔融形成的花岗质岩浆(可能伴有少量幔源物质混入)及其分异出的岩浆热液萃取基底成矿物质后,成矿流体迁移至NE向构造发生矿质沉淀的产物。

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