

# 德兴斑岩铜矿田黄铁矿 Re-Os 同位素定年 及其地质意义 \*

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**摘要** 德兴斑岩铜矿田是中国东部最大的斑岩铜矿系统, 一直以来, 德兴铜矿的成矿过程和成因机制都是矿床学家关注的热点问题。前人在德兴铜矿的围岩蚀变、矿化期次、流体特征、成岩年代等方面取得了较为一致的认识, 但其成矿时代仍争议较大。文章首次对矿田内与黄铜矿密切共生的黄铁矿进行了 Re-Os 同位素定年, 10 件样品中的 4 件  $w(\text{Re})$  极低, 模式年龄变化范围大, 且误差较大, 故不对其进行讨论; 其余 6 件的  $w(\text{Re})$  为  $10.58 \times 10^{-9} \sim 102.59 \times 10^{-9}$ , 普通  $w(\text{Os})$  较低 ( $0.0054 \times 10^{-9} \sim 0.0113 \times 10^{-9}$ ),  $w(^{187}\text{Os})$  较高 ( $0.019 \times 10^{-9} \sim 0.177 \times 10^{-9}$ ),  $\text{Re}/\text{Os}$  比值较高 (4406~73 422), 为低含量高放射性 Os 成因硫化物。研究获得黄铁矿 Re-Os 同位素加权平均模式年龄为  $(165.3 \pm 2.3)\text{Ma}$  ( $\text{MSWD}=1.04$ )。因此, 德兴铜矿的铜成矿年龄为中侏罗世, 综合前人资料, 德兴铜矿最可能形成于古太平洋板块俯冲的远程效应影响下的陆内伸展的地质背景, 为陆内环境斑岩铜矿。

**关键词** 地球化学; 黄铁矿; Re-Os 同位素; 成矿时代; 德兴斑岩铜矿田

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## Re-Os isotopic dating of pyrite from Dexing porphyry copper orefield and its geological significance

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### Abstract

The Dexing porphyry copper orefield is the largest porphyry Cu system in eastern China. The metallogenetic process and genetic mechanism of Dexing copper mine have always been a hot issue which have aroused much interest among geologists. Previous studies have reached a consensus on the alteration, mineralogic stages, characteristic of fluids and petrogenetic age. However, its metallogenetic age remains controversial. In order to further determine its metallogenetic age, the authors first carried out the Re-Os isotopic dating of pyrite. Four pyrite samples were rejected for their extremely low content of Re, wide range of model ages and large errors. The remaining six pyrite samples belong to low-level highly radiogenic (LLHR) sulfides, characterized by Re content between  $10.58 \times 10^{-9}$  and  $102.59 \times 10^{-9}$ , low common Os content ( $0.0054 \times 10^{-9} \sim 0.0113 \times 10^{-9}$ ), relatively high content of  $^{187}\text{Os}$  ( $0.019 \times 10^{-9} \sim 0.177 \times 10^{-9}$ ) and Re/Os ratios (4406~73 422). The average Re-Os model age is  $(165.3 \pm 2.3)$

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Ma (MSWD=1.04), indicating that the porphyry style Cu mineralization occurred during the Jurassic period. The Dexing copper mine was probably formed in the geological background of intracontinental extension.

**Key words:** geochemistry, pyrite, Re-Os isotope, metallogenetic age, Dexing porphyry copper orefield

近几十年来,Re-Os同位素逐渐成为探究金属硫化物矿床成矿年代和成矿物质来源的直接和有效手段(蒋少涌等,2000;杜安道等,2009)。因辉钼矿具有较高的Re含量,而几乎不含普通Os,其已成为目前应用最广泛的Re-Os定年矿物(谢桂青等,2006;丰成友等,2012;范羽等,2014;李海立等,2016;Li et al., 2017; Ni et al., 2017)。相比于辉钼矿,黄铁矿在金属硫化物矿床中更为普遍,但是由于其含有较低的Re、Os含量(通常 $10^{-9}$ ~ $10^{-12}$ 级),通常不容易得到较好的等时线年龄(蒋少涌等,2000)。黄铁矿Re-Os同位素定年在中国主要成功应用于金矿(石桂勇等,2012;赵晓波等,2014; Ding et al., 2016; 张朋等,2016)、铅锌矿(唐永永等,2013; Li et al., 2016)、铁矿(Huang et al., 2013; 邵建波等,2014; 刘明军等,2014)、铜矿及锡多金属矿等(梁婷等,2009)。对于铜矿而言,黄铁矿Re-Os同位素定年的成功案例主要包括海底热水沉积-热液叠加改造型铜矿(郭维民等,2011;陈雷等,2013)、铜镍硫化物矿床(Wang M F et al., 2015; 段士刚等,2017)、矽卡岩型铜矿(陈红瑾等,2011; Ying et al., 2014)和斑岩型铜矿(Zhang et al., 2016)等。

德兴斑岩铜矿田是华南最大的斑岩铜矿系统,自1958年建矿以来就有众多地质学家对其进行研究。在德兴铜矿的成岩时代方面,近期研究者对花岗闪长斑岩开展高精度SHRIMP或LA-ICP-MS锆石U-Pb测年,得到了较为统一的年代,约171 Ma (Wang et al., 2006; Liu et al., 2012; Zhou et al., 2012a; Li et al., 2013; Wang G G et al., 2015)。此外,德兴铜矿还广泛发育石英闪长玢岩等脉岩,Li等(2013)和李利等(2018)利用锆石SHRIMP或LA-ICP-MS U-Pb法得到朱砂红、富家坞矿床闪长玢岩的年龄为168~166 Ma,略晚于花岗闪长斑岩的侵位时间。

德兴铜矿的成矿时代还存在争议。近年来,一些研究者根据德兴铜矿田的辉钼矿Re-Os年龄,认为德兴铜矿的成矿年代约为171 Ma (Guo et al., 2012; Zhou et al., 2012a; Li et al., 2017)。也有研究者根据与铜矿化密切相关的金红石和磷灰石的ID-TIMS U-Pb定年方法,获得铜矿化年代为165 Ma (Li

et al., 2013)。Li等(2013)结合辉钼矿Re-Os年代学资料提出德兴铜矿存在钼矿化(171 Ma)和铜矿化(165 Ma)两期热液事件。此外,Zhou等(2012b)通过对大脉型含矿石英脉中热液锆石SHRIMP U-Pb定年(~104.3 Ma),认为德兴斑岩铜矿田还存在另一期早白垩世的成矿事件。为进一步确定德兴斑岩铜矿田的铜矿化时代,本文首次对德兴斑岩铜矿田内的与黄铜矿密切共生的黄铁矿进行Re-Os同位素定年,并探讨其地质意义。

## 1 区域地质背景

德兴斑岩铜矿田位于扬子地块东南缘,江(山)—绍(兴)断裂带北西侧,江南造山带东段(图1a)。区内地层为新元古界双桥山群(图1b),由浅海沉积的泥砂岩和火山碎屑岩夹凝灰熔岩组成,这些碎屑岩经历了低级区域绿片岩相变质作用之后形成了一套厚2678~5472 m的千枚岩(Wang et al., 2008)。上覆地层为中侏罗统鹅湖岭组(图1b),其出露在德兴地区西南部NE向的孔家-银山火山岩盆地中,LA-ICP-MS锆石U-Pb定年表明此次火山事件发生在中侏罗世(176~166 Ma)(Wang et al., 2012)。赣东北断裂为区域内的一级断裂(图1b),东北向延伸,长度超过200 km。蛇绿混杂岩(1.0 Ga)沿该断裂分布,代表了新元古代扬子板块和海洋岛弧的次级缝合带(Li et al., 1997)。另外,区内北西部为北东向展布的乐安江断裂,中部为泗洲庙向斜(图1b)。

德兴地区的岩浆侵入事件发生在晚中生代(侏罗纪—早白垩纪)(图1b)。德兴花岗闪长斑岩(锆石U-Pb年龄为170 Ma)和银山英安岩到流纹质次火山岩(锆石U-Pb年龄为170~166 Ma)均是侏罗纪岩浆事件的产物(Wang et al., 2012; 2015)。区内东南部的大型A型大茅山花岗岩为白垩纪的产物(锆石U-Pb年龄为126~122 Ma; Jiang et al., 2011)。

## 2 矿田地质

德兴斑岩铜矿田由朱砂红、铜厂和富家坞3个矿床组成,呈北西西方向展布(图2)。矿田内出露地

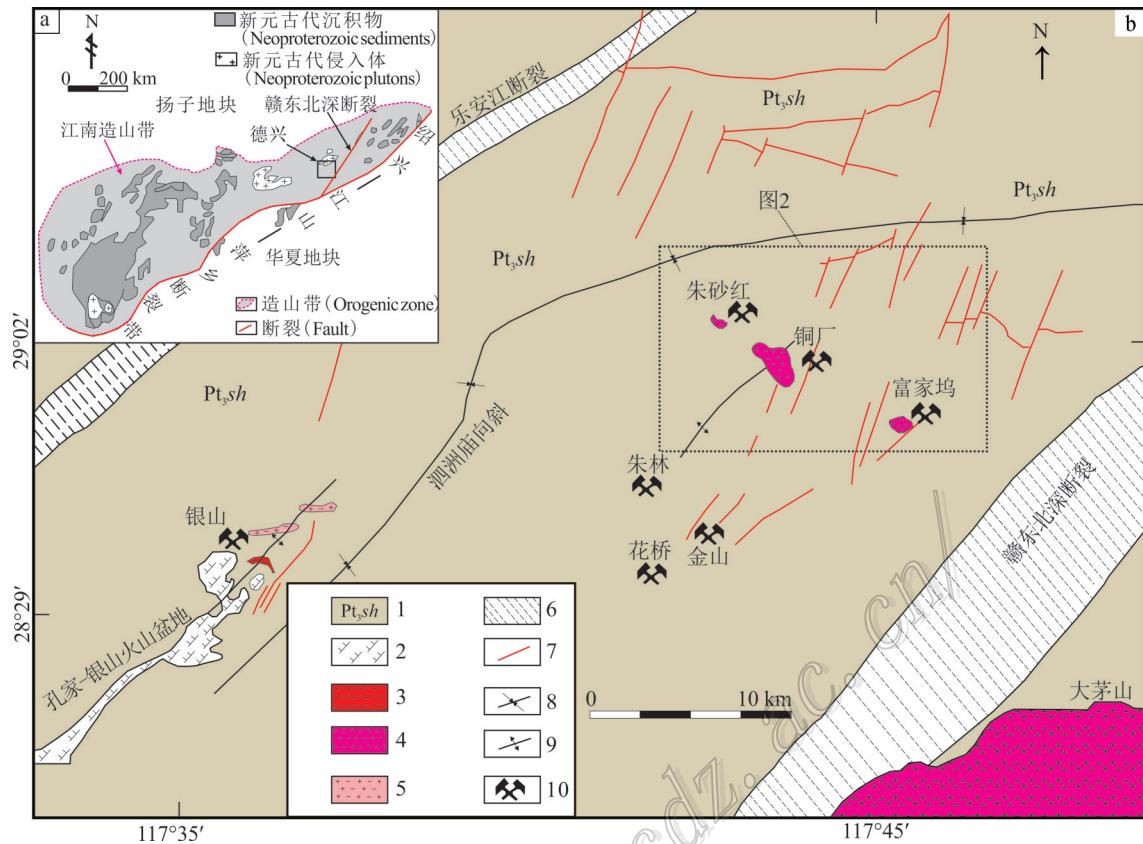


图1 华南江南造山带地质简图(a, 据Wang et al., 2008修改)和德兴地区地质简图  
(b, 据Wang G G et al., 2015修改)

1—新元古代双桥山群浅变质岩;2—中侏罗世鹅湖岭组火山岩;3—中侏罗世安山玢岩;4—中侏罗世花岗闪长斑岩;5—中侏罗世石英斑岩;6—深断裂;7—断层;8—向斜;9—背斜;10—矿床

Fig. 1 Simplified geological map of the Jiangnan orogen in South China (a, modified after Wang et al., 2008)

and sketch map of the Dexing ore district (b, modified after Wang G G et al., 2015)

1—Neoproterozoic Shuangqiaoshan Group epi-metamorphic phyllite; 2—Middle Jurassic Ehuling Formation volcanic rock;  
3—Middle Jurassic dacitic porphyries; 4—Middle Jurassic granodiorite porphyries; 5—Middle Jurassic quartz  
porphyries; 6—Deep fault zone; 7—Fault; 8—Synclinorium; 9—Anticline; 10—Deposit

层为新元古界双桥山群浅变质岩。矿田北侧的泗州庙向斜是矿田内的主要褶皱构造,另外还发育北东向次级褶皱,如西源岭背斜、官帽山向斜、富家坞鼻状构造等。矿田中发育多个方向的断裂构造,矿田的内挤压断裂包括东西向、北东向和北北东向的断裂系统,这些断裂都是与同方向的褶皱构造相伴出现,另外还可见北西西向的横张断裂(图2)。

矿田内的3个矿床均主要与燕山期花岗闪长斑岩有关,单个岩体均向北西深部倾伏,呈3个大小不同的似筒状岩株。3个主岩体与围岩均呈接触侵入,接触界线清楚,呈突变关系。3个矿床的铜钼矿体均产于含矿花岗闪长斑岩岩体顶部和上部的内外接触带,岩体的中心及深部,一般未形成工业矿体,斑岩

中的铜矿体小于围岩中的铜矿体,其空间形态呈倾向北西的空心筒状体,在水平切面上呈“环形”,在纵剖面上呈倾斜的“梭形”。从北西部的朱砂红矿床到南东部的富家坞矿床,矿体倾角由陡变缓,矿体形态由复杂到简单,矿石铜钼品位由贫变富,矿床剥蚀程度由浅变为中等。矿田内3个矿床具有相似的蚀变类型和分带型式,与典型斑岩铜矿的蚀变类型一致,即钾化蚀变、黄铁绢英岩化蚀变和青磐岩化蚀变(图2)。

### 3 样品特征及测试方法

本次用于Re-Os同位素年龄测定的10件黄铁矿

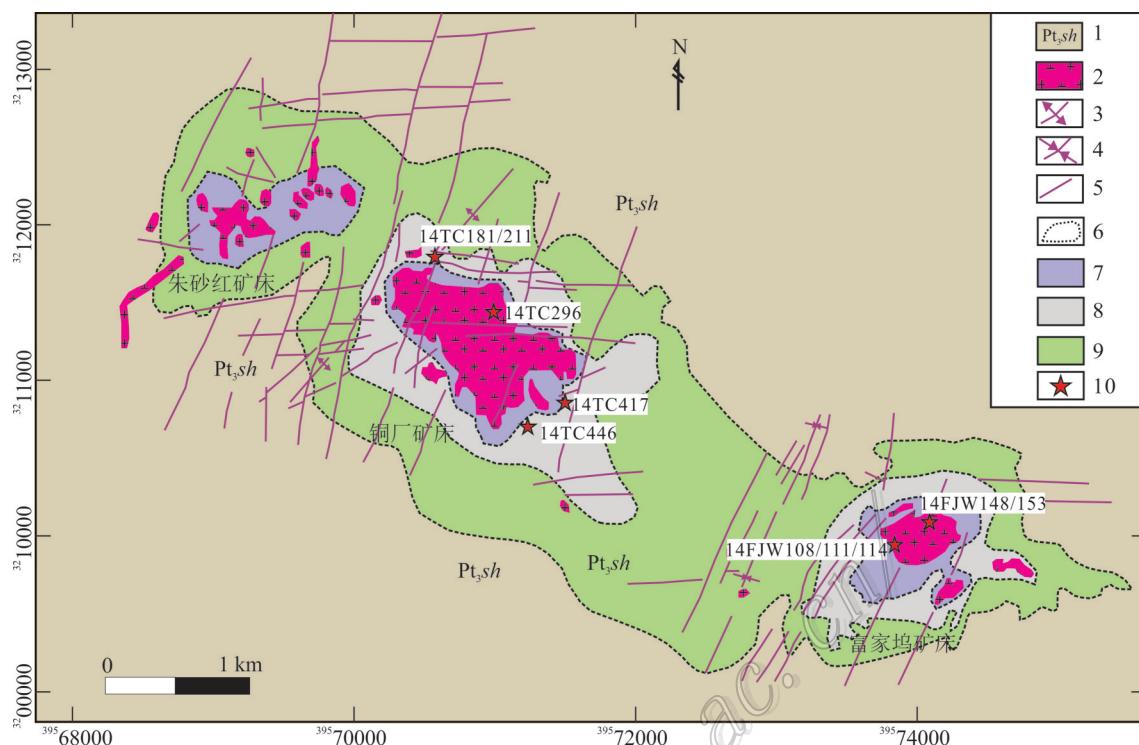


图2 德兴铜矿田地质简图和样品位置(据 Wang G G et al., 2015 修改)

1—新元古代双桥山群浅变质岩;2—中侏罗世花岗闪长斑岩;3—背斜;4—向斜;5—断层;6—蚀变界线;  
7—钾化蚀变;8—黄铁绢英岩化蚀变;9—青磐岩化蚀变;10—采样位置

Fig. 2 Geological map of Dexing porphyry copper orefield and sample location  
(modified after Wang G G et al., 2015)

1—Neoproterozoic Shuangqiaoshan Group epi-metamorphic phyllite; 2—Middle Jurassic granodiorite porphyry;  
3—Anticline; 4—Syncline; 5—Fault; 6—Alteration boundary; 7—Potassic alteration;  
8—Phyllitic alteration; 9—Propylitic alteration; 10—Sample location

样品,其中5件来自铜厂矿床,其余的来自富家坞矿床,采样位置见图2。样品均为矿石,主要呈细脉浸染状产出,发育半自形-他形粒状结构、粒间充填交代结构、裂隙充填交代结构等,矿石矿物主要为黄铁矿、黄铜矿、辉钼矿等,黄铜矿与黄铁矿紧密共生,同时可见辉钼矿与黄铜矿、黄铁矿共生,脉石矿物主要为石英、绢云母等(图3a~d)。样品经人工粉碎后,按重力和磁选方法分选出黄铁矿,最后在双目镜下挑选(纯度>99%)。

样品分析和测试工作在中国科学院地球化学研究所矿床地球化学国家重点实验室完成,实验仪器为ELAN DRC-e型电感耦合等离子体质谱仪(ICP-MS),仪器灵敏度大于4000 cps/ $10^{-9}$   $^{115}\text{In}$ , 相对标准偏差(RSD%)小于3%。实验中所用的HCl和HNO<sub>3</sub>通过亚沸蒸馏提纯,实验用水为Millipore 18 MΩ·cm,  $^{187}\text{Re}$  和  $^{190}\text{Os}$  同位素稀释剂分别稀释至

$10 \times 10^{-9}$  备用,卡洛斯(Carius)管约230 mm长,内径26 mm,外径30 mm,内部容积大约120 mL,实验前用60%的王水加热煮5 h,蒸馏水清洗干净后加热至560°C约8 h退火。

实验步骤简述如下:准确称取3~5 g黄铁矿样品置于Carius管中,缓慢加入HNO<sub>3</sub>以分解黄铁矿,吸收管将产生的气体和少量Os用2.5 mL 10 mol/L的HCl在冰水浴中吸收,反应过程中需持续通入净化的空气,以免Carius管冷却使其内压降低而吸收Os的HCl回流。反应完全后,将吸收Os的HCl转移到Carius管中,并用4 mL HCl分2次清洗Carius管,然后加入适量的Re和Os稀释剂,加入15 mL HNO<sub>3</sub>,密封Carius管,将其置入不锈钢套中在200°C下加热15 h,然后取出Carius管并置于冰箱的冷却室中冷却2 h,打开Carius管,将进气的特氟龙(Teflon)管移至Carius管底部,在沸腾的水浴中加热蒸馏,在蒸馏过

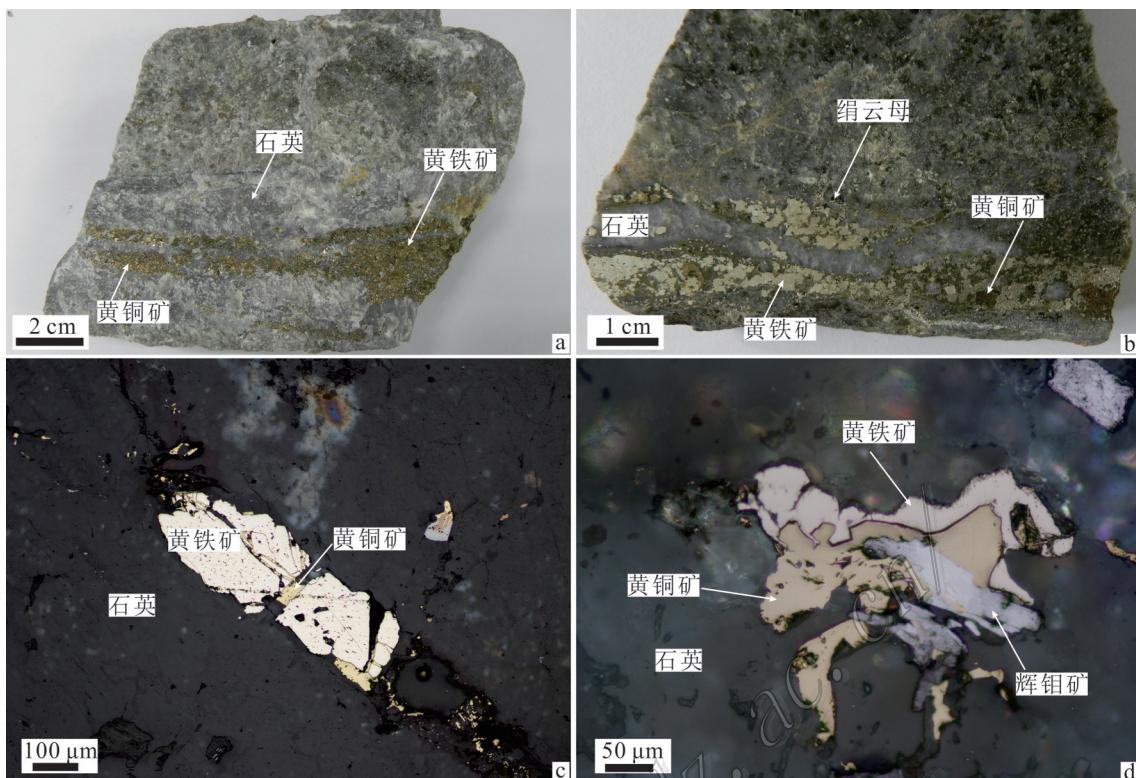


图3 德兴斑岩铜矿田黄铁矿Re-Os同位素样品典型手标本照片(a、b)和显微照片(c、d)

Fig. 3 Representative specimen photos (a, b) and reflected light microphotographs (c, d)  
for pyrite Re-Os isotope sample in the Dexing porphyry copper orefield

程中缓慢加入约5 mL H<sub>2</sub>O<sub>2</sub>使Os蒸馏完全, Os用净化的空气带出, 并用1.5 mL H<sub>2</sub>O在冰水浴中吸收, 将吸收液保存在冰箱冷冻室, 备测。将蒸馏Os剩下的溶液转移至125 mL Teflon烧杯中蒸干, 加入3 mL HCl再蒸干, 用25 mL 2 mol/L的HCl溶解残渣, 用阴离子交换树脂(AG1-X8 200~400目)分离Re。实验原理及详细步骤参阅杜安道等(2009)、漆亮等(2009)和Qi等(2010)。

#### 4 测试结果

10件黄铁矿样品的Re-Os同位素分析结果见表1。其中铜厂矿床有4件样品的Re含量极低( $<3 \times 10^{-9}$ ), 从而导致其模式年龄变化范围大(2192.6~240.0 Ma), 且误差较大(>10 Ma), 无法形成等时线谱和图, 因而不对其进行讨论。其余6件测试样品的 $w(\text{Re})$ 为 $10.58 \times 10^{-9} \sim 102.59 \times 10^{-9}$ , 普通 $w(\text{Os})$ 较低( $0.0054 \times 10^{-9} \sim 0.0113 \times 10^{-9}$ ),  $w(^{187}\text{Os})$ 较高( $0.019 \times 10^{-9} \sim 0.177 \times 10^{-9}$ ), Re/Os比值较高(4406~73 422), 模

式年龄集中在173.5~162.3 Ma, 普通Os相对于放射性Os可忽略不计, 与Stein等(2000)提出的低含量高放射成因Os(LLHR)的硫化物类似。由于此类样品中的Os主要以放射性<sup>187</sup>Os为主, 普通Os含量很低, 导致普通Os很难准确测定, 造成<sup>187</sup>Re/<sup>188</sup>Os与<sup>187</sup>Os/<sup>188</sup>Os之间的相关误差较大, 因此, 选用以<sup>187</sup>Re为横坐标, 以<sup>187</sup>Os为纵坐标的等时线投图, 利用ISOPLOT程序得到<sup>187</sup>Re-<sup>187</sup>Os等时线年龄为(162.7±6.9) Ma(MSWD=0.48)(图4a), 加权平均模式年龄为(165.3±2.3) Ma(MSWD=1.04)(图4b)。

#### 5 讨论

##### 5.1 德兴铜矿黄铁矿Re-Os同位素年龄

Re、Os元素具有亲铁、亲硫及耐熔的特性, 在自然界它们主要富集于金属和硫化物相中, 是直接约束金属矿床成矿时代的有力工具。本次选取的黄铁矿与黄铜矿密切共生, 因此, 其Re-Os同位素约束的成矿时代可以代表铜矿化时代。

表1 德兴斑岩铜矿田黄铁矿 Re-Os 同位素结果

Table 1 Re-Os isotopic data of pyrite from the Dexing porphyry copper orefield

| 样品号                | $w(\text{Re})/10^{-9}$ |           | $w(\text{普 Os})/10^{-9}$ |           | $w(^{187}\text{Re})/10^{-9}$ |           | $w(^{187}\text{Os})/10^{-9}$ |           | 模式年龄/Ma |           | Re/Os<br>比值 |
|--------------------|------------------------|-----------|--------------------------|-----------|------------------------------|-----------|------------------------------|-----------|---------|-----------|-------------|
|                    | 测定值                    | $1\sigma$ | 测定值                      | $1\sigma$ | 测定值                          | $1\sigma$ | 测定值                          | $1\sigma$ | 测定值     | $1\sigma$ |             |
| 等时线谱和样品            |                        |           |                          |           |                              |           |                              |           |         |           |             |
| 14TC181            | 10.58                  | 0.34      | 0.0113                   | 0.0017    | 6.62                         | 0.21      | 0.019                        | 0.001     | 173.5   | 8.0       | 4406        |
| 14FJW153           | 30.92                  | 0.42      | 0.0059                   | 0.0021    | 19.35                        | 0.27      | 0.056                        | 0.001     | 172.2   | 4.4       | 24543       |
| 14FJW148           | 24.55                  | 1.87      | 0.0054                   | 0.0006    | 15.37                        | 1.17      | 0.042                        | 0.001     | 165.0   | 3.0       | 21439       |
| 14FJW111           | 83.75                  | 2.66      | 0.0058                   | 0.0010    | 52.43                        | 1.67      | 0.145                        | 0.002     | 165.7   | 1.8       | 68070       |
| 14FJW114           | 57.55                  | 0.92      | 0.0058                   | 0.0009    | 36.02                        | 0.57      | 0.098                        | 0.001     | 162.3   | 2.4       | 46850       |
| 14FJW108           | 102.59                 | 2.84      | 0.0066                   | 0.0014    | 64.22                        | 1.78      | 0.177                        | 0.004     | 164.9   | 3.6       | 73422       |
| $w(\text{Re})$ 低样品 |                        |           |                          |           |                              |           |                              |           |         |           |             |
| 14TC211            | 0.92                   | 0.09      | 0.0117                   | 0.0056    | 0.58                         | 0.06      | 0.004                        | 0.000     | 380.9   | 26.7      | 370         |
| 14TC296            | 1.05                   | 0.02      | 0.0051                   | 0.0008    | 0.66                         | 0.02      | 0.003                        | 0.000     | 240.0   | 12.3      | 971         |
| 14TC417            | 2.08                   | 0.06      | 1.0769                   | 0.0018    | 1.30                         | 0.04      | 0.007                        | 0.001     | 300.4   | 20.7      | 1288        |
| 14TC446            | 0.05                   | 0.03      | 0.5007                   | 0.0071    | 0.03                         | 0.21      | 0.001                        | 0.000     | 2192.6  | 278.8     | 32          |

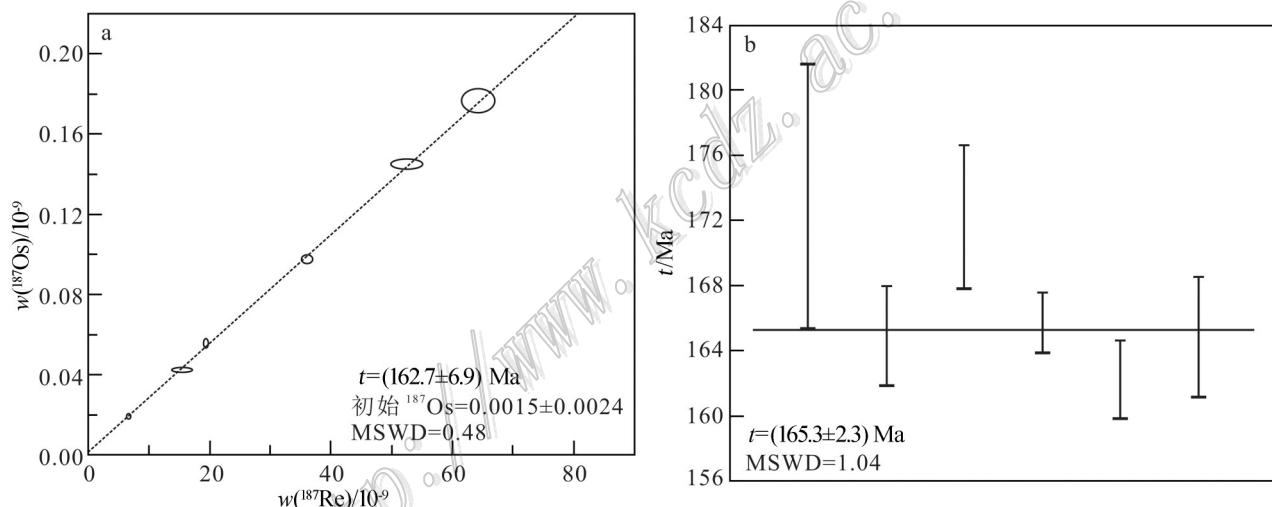


图4 德兴斑岩铜矿田黄铁矿 Re-Os 同位素等时线年龄(a)和加权平均模式年龄(b)

Fig. 4 Re-Os isochron age (a) and Re-Os weighted average model ages (b) of pyrite from the Dexing porphyry copper orefield

德兴斑岩铜矿田成矿年代学研究始终是中国地质学者的研究热点,长期以来,在该区获得了大量的成岩成矿年龄数据。朱训等(1983)通过全岩和矿物的K-Ar、Rb-Sr同位素得到花岗闪长斑岩的等时线年龄为172~112 Ma,石英闪长玢岩的等时线年龄为193~96 Ma。由于全岩或矿物的K-Ar和Rb-Sr封闭温度较低,后期低温的热液蚀变会导致岩体的Rb-Sr和K-Ar同位素不封闭而导致Rb-Sr和K-Ar同位素年龄不能准确代表岩浆结晶年龄。而锆石性质稳定,不易受到蚀变的影响,而且其U-Pb体系封闭温

度很高(900~1100°C),因此,锆石U-Pb同位素方法可以获得准确的岩体侵位年龄(Cliff, 1985; Lee et al., 1997)。Wang等(2006)、Liu等(2012)、Zhou等(2012a)和Wang G G等(2015)通过锆石SHRIMP或LA-ICP-MS U-Pb定年得到花岗闪长斑岩的年代为172~170 Ma。Li等(2013)和李利等(2018)通过锆石SHRIMP或LA-ICP-MS U-Pb定年得到闪长玢岩的年代为168~166 Ma。故德兴斑岩铜矿田内的岩浆活动发生在侏罗纪(172~166 Ma)。

Guo等(2012)、Zhou等(2012a)、曲焕春等

(2015) 和 Li 等(2017) 通过辉钼矿 Re-Os 同位素开展成矿年代研究, 推断斑岩铜矿成矿时代为中侏罗世, 约 171 Ma。尽管辉钼矿 Re-Os 年代学定年准确性高, 可信度高, 但是部分学者提出德兴铜矿的钼、铜可能不是同期形成的产物(Li et al., 2013)。Li 等(2013) 通过铜矿化相关的磷灰石和金红石 ID-TIMS U-Pb 定年, 推断铜矿化略晚于钼矿化, 斑岩铜矿化时代为 165 Ma。此外, Zhou 等(2012b) 通过热液锆石 U-Pb 定年得到德兴斑岩铜矿田内可能存在白垩纪的铜成矿事件(106~100 Ma)。

铜厂和富家坞矿床 6 个较高 Re 含量的黄铁矿样品中有 2 个样品的模式年龄为 172.2 Ma 和 173.5 Ma, 其余 4 个样品模式年龄比较年轻, 为 165.7~162.3 Ma。前人报道的德兴地区花岗闪长斑岩和辉钼矿 Re-Os 年龄均约为 171 Ma, 与本次 2 件黄铁矿样品的模式年龄一致。国际上也有报道指出, 斑岩型矿床的成岩成矿时差常很小, 通常小于 1~2 My (Von Quadt et al., 2011)。本次研究的结果也暗示了德兴铜矿铜、钼矿化为同期产物。

对于模式年龄年轻的黄铁矿样品, 存在 2 种可能性: 第一种为分析测试导致的误差, 这主要是由于黄铁矿样品中 Re 含量远低于辉钼矿中 Re 含量, 因此, 黄铁矿样品中 Re-Os 体系微小的同位素组成测定误差可能会产生较大的模式年龄误差; 第二种可能性为分析对象为黄铁矿可能经历了后期热液蚀变的影响, 即早期黄铁矿(~171 Ma)受到后期热液蚀变的改造导致, 后期的热液事件年代时间尚不能确定, 可能为~165 Ma 闪长玢岩对应的热液蚀变(金红石和磷灰石 TIMS U-Pb 年龄为~165 Ma)或者为 104.3 Ma 热液蚀变造成。综上所述, 本文倾向于解释黄铁矿 Re-Os 定年的数据为德兴铜矿主体铜成矿时间为 171 Ma, 后期热液蚀变对成矿贡献大小需要进一步研究。

## 5.2 成矿地质背景

华南中生代构造动力学演化模型存在的观点主要可分为 3 类: ① 碰撞模型(Hsü et al., 1988); ② 地幔柱模型(谢窦克等, 1996; 谢桂青等, 2001); ③ 古太平洋或 Izanagi 板块俯冲模型(Zhou et al., 2006; Li et al., 2007; Sun et al., 2007)。目前被相关学者普遍接受的为古太平洋板块俯冲模型(毛景文等, 2011; Jiang et al., 2015; Wang G G et al., 2015)。Wang 等(2016)认为中侏罗世—早白垩世俯冲于华南陆块之下的更可能为伊泽纳崎板块(或古太平洋

板块)(Engebretson, 1985), 并且俯冲板块漂移方向为 NW 向或 NNW 向(Goldfarb et al., 2007; Seton et al., 2012)。

本文从成矿年代方面, 进一步确定了德兴铜矿的成矿时代为中侏罗世。中侏罗世伊泽纳崎板块北西向俯冲到华南板块之下, 而此时华南活动大陆边缘弧岩浆作用微弱, 可能是由于伊泽纳崎板块未俯冲至~110 km 的深度(Tatsumi et al., 1995)。但是, 压力可能已经传到了华南内陆, 活化了早先存在的深断裂(如赣东北深断裂)(Zhou et al., 2006; He et al., 2010; Wang et al., 2012)。深断裂的活化导致了陆内拉伸, 这可能使得华南内陆的南岭地区形成 A 型花岗岩、双峰式火山岩和碱性玄武岩(Chen et al., 2005; 朱金初等, 2008; He et al., 2010)。靠近赣东北断裂, 德兴斑岩铜矿田内的花岗闪长斑岩、石英闪长玢岩脉岩, 形成于陆内拉伸的构造背景。在此背景下, 新元古代江南造山运动形成的富铜初始岩石圈(Wang G G et al., 2015)拆沉或直接部分熔融形成了德兴斑岩铜矿的高镁埃达克岩浆(即花岗闪长斑岩岩浆), 岩浆上涌, 形成德兴斑岩铜矿。

## 6 结 论

(1) 德兴斑岩铜矿田内黄铁矿为低含量高放射性成因 Os 硫化物, 黄铁矿 Re-Os 加权平均模式年龄为  $(165.3 \pm 2.3)$  Ma, 说明其成矿作用发生在中侏罗世。

(2) 德兴铜矿最可能形成于古太平洋板块俯冲的远程效应影响下的陆内伸展地质背景, 为陆内环境的斑岩型铜矿。

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